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we extracted the daily surface temperature dataset from two cmip5 models: cesm4 and bcc-csm1.1, which are the only 2 cmip5 models used to construct the corresponding observation data. the data are derived from the simulations performed between 1993 and 2012. first, the model data are scaled to match the era-20c observations at the time of analysis. the final differences between the two reanalysis datasets and the observational data are driven by the grid sampling, which is different in the reanalysis data. second, we calculate a 30-year running annual mean temperature dataset. then, we create a standardized bias corrected version of both 30-yr running mean temperature trends over the entire us (fig. 23a) and over the eastern us (fig. 23b). all dates used in the trends are from april 1 to june 30 except for december and january, which are excluded because there are few observations over december and january in the area. as described in palmer et al. (2013), we used the output data from 10 cmip5 climate models, with the following models having common initial and boundary conditions: bcc-csm1.1, cesm1-bgc, cesm1-cam5, cesm1-ccsm3, cesm1-cccam2, cesm1-cam5-osaka, hadgem1-cc, hadgem2-es, inmcm4-acme2, and ipsl-cm5a-lr. these 10 models all use atmospheric general circulation model (gcms) and include either an atmosphere-only or coupled ocean-atmosphere model and a control configuration with both atmospheric and oceanic variables as input. the differences in all variables between the 10 gcms are small, the smallest difference in air temperature over land is 0.27 k. the fact that gcms show relatively little difference in simulation is consistent with the fact that we have a large spread in observations. therefore, the variations in the bias corrected 30-yr mean trends at the annual scale over the entire us and over the eastern us can be attributed to differences in the models. figure 24 shows the results for the air temperature trends over the us and eastern us in the 10 models. all models have a trend with a positive sign over the us, and a negative over the eastern us. results from both the 10 models and the data in palmer et al. (2013) show a significant positive trend over the us with the cmip5 models showing larger positive trends than those from the observations. the differences in the 10 models on the eastern us are, however, small, with the cmip5 models having a near zero trend, while the results from the observation show a significant negative trend over the eastern us.

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we examine the mean global pattern of temperature, precipitation, sea-level pressure, and the mjo in the cmip5 ensemble. figure 20 shows the observations and the ensemble mean of the ensemble mean n-member solutions m= n (eτm n). the overall pattern of the global average temperature, precipitation, and mjo is similar to the observations, although the models predict weaker mjo events, particularly over the tropical pacific (location of the thermocline layer of the mjo) than the observations. the mjo-related changes in sea-level pressure show a good correlation between observations and the models. the global mean sea-level pressure is lower in the models than in observations in the tropical pacific. the mjo related patterns of temperature and rainfall over south america and western north america appear to be better reproduced in the cmip5 models than in the observations. however, they both fail to reproduce the strong mjo intensity over the eastern u.s. and southeast asia. figure 21 shows the observed and predicted anomalies of five variables in the area around the united states. the relative humidity is predicted to be higher over north america than in observations, whereas the mean sea-level pressure is higher in observations than in the model mean. on the other hand, we find a good agreement for temperature and precipitation. the different phase of the pdo in each model is shown in fig. 21a and clearly defines the different pattern over north america for the models. the best simulation of the pdo is presented by the model simulations of the warm phase, which is associated with a positive phase of the pdo. in contrast, the cold phase, characterized by a negative phase of the pdo, is somewhat better reproduced by the cold-phase simulations of the models. the hadgem2-es, norwayesm, and canesm2-es models are associated with the cold phase (pdo (fig. 21b). the cold-phase simulations replicate relatively well the cold phase of the pdo, although the wintertime precipitation (figs. 21d,f) do not suggest a relationship with the pdo as captured by the observations. the wintertime precipitation (figs. 5ec8ef588b

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